

University of Maryland Baltimore County - UMBC

Phys650 - Special Topics in Experimental Atmospheric Physics (Spring 2009)

J. Vanderlei Martins and Manfredo H. Tabacniks

<http://userpages.umbc.edu/~martins/PHYS650/>

Class 5

- In Depth Discussion and Continuation of the Course Project
- Introduction to Radiative Transfer: Theory and Codes
- Introduction to radiometry
- Hands on Experiment: Calibration and Characterization of the Sunphotometer

In depth discussion and continuation of the course project:

- Introduction to Radiative Transfer and Radiative Transfer codes
- Introduction to Radiometric Analysis:
 - How many photons arrive at the surface of your detector?
 - What is the uncertainty of your measurement?
- Discussion on the FOV of the Photometer
 - What is the FOV of your system?
 - What is the FOV of the sun on the sky?
 - What is the associated error from the difference between the above FOVs?
- Discussion on the Calibration of the Photometer
 - What is the required accuracy for sunphotometry and what method can meet this requirement?
 - Calibration with Langley Plots
 - Calibration in Integrating Spheres

After calibration, you can start measuring the aerosol optical depth at some location and compare it with Aeronet measurements, Lidar, or satellite retrievals. Notice that you cannot touch/modify your photometer after calibration.

- Discussion on measurements of sky radiance, moon photometry, etc.
 - How can you convert your sunphotometer to a sky radiometer?
 - How many photons do you get from the sky at the surface of your detector?
 - Can you convert your sunphotometer to a moon, or a star photometer?
 - How many photons do you need arriving at the surface of your detector in order to make an accurate measurement in each case?
 - How can you increase the number of photons arriving at the surface of your detector

Experiments

12:00-12:30 - Hands on Experiment:

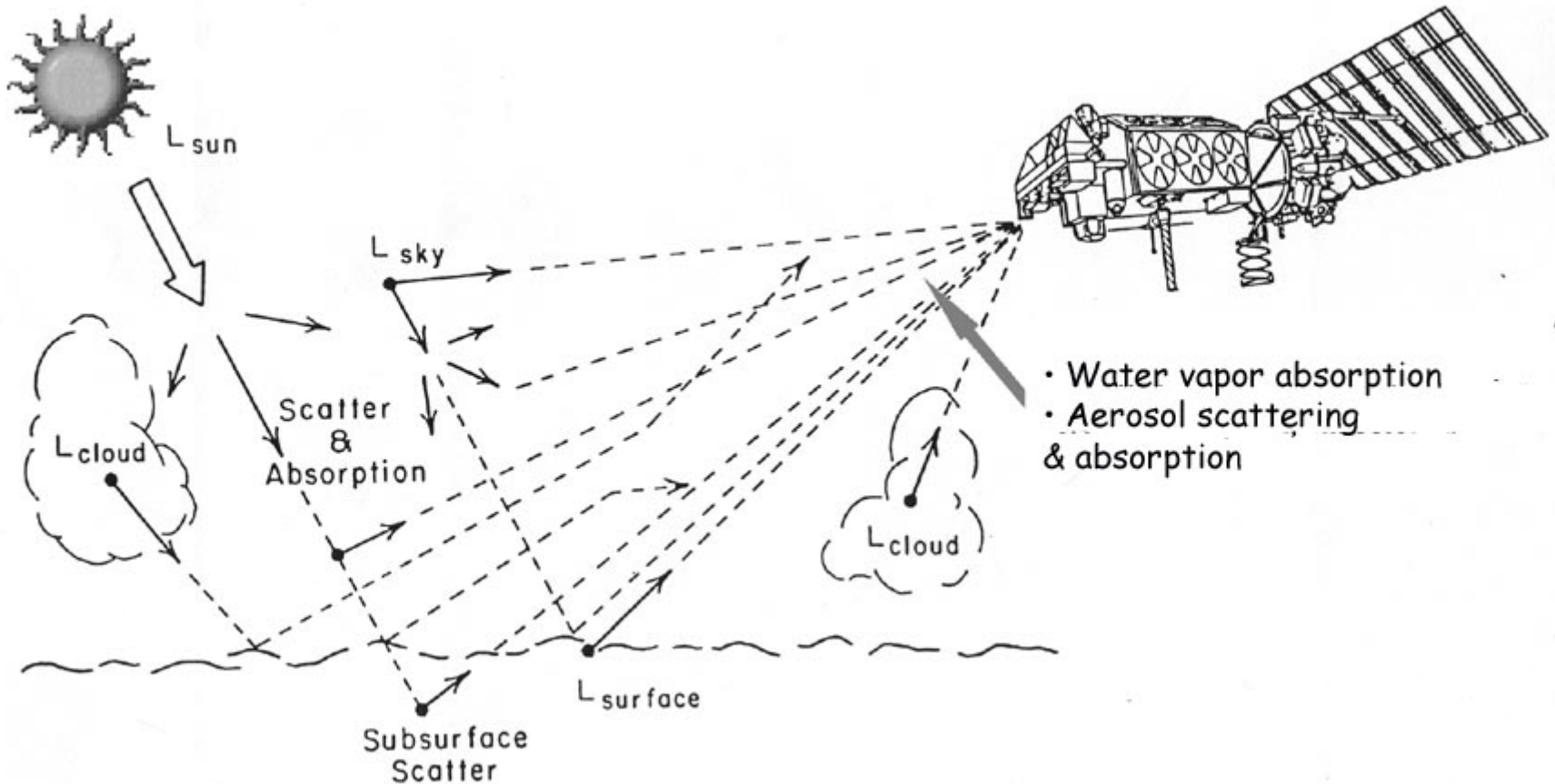
- Assemble a "Tiger team" to design your instrument and attack the questions above. Each team should address all the questions above theoretically and experimentally. Try to explore the expertise of the team. You should understand the whole experiment but, if you are better in theory or in computation, go for it and help the team to achieve its goals!!!
- Visit to the Integration sphere facility. Experiments with the sphere should be planned in advance. The sphere can only be used under supervision.

12:30- 3:00 PM - Open Lab:

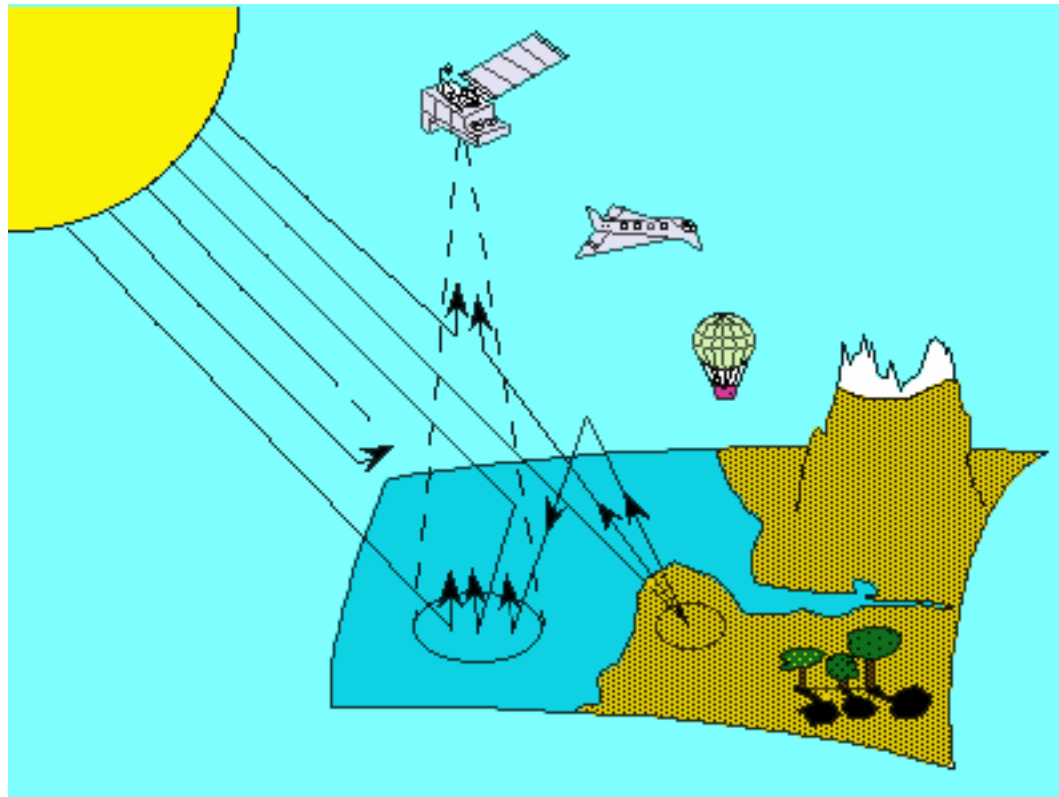
Tiger Teams shall design and start their experimental setups to measure the properties of their photometers. Measurements will include:

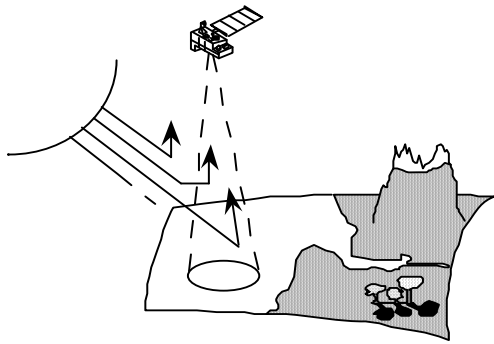
- Measurements of the LED FOV (emission and detection)
- Alignment of the LED detection via emission
 - I have noticed a dark spot at the center of some LEDs!!! Double check if it have a negative effect on the detection!!!
- Radiometric calibration at the integrating sphere

Radiance sources and sinks affecting visible and IR wavebands

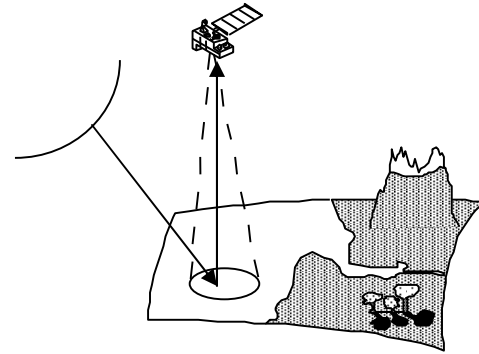


Solar Energy Paths

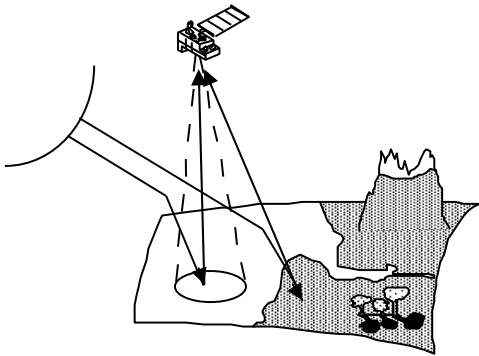




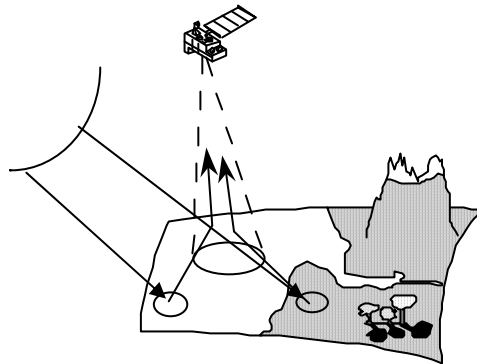
atmospheric contribution



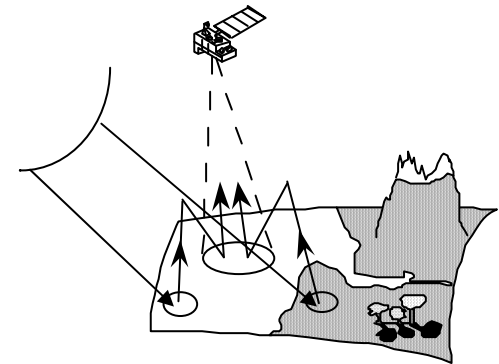
direct + direct



diffuse + direct

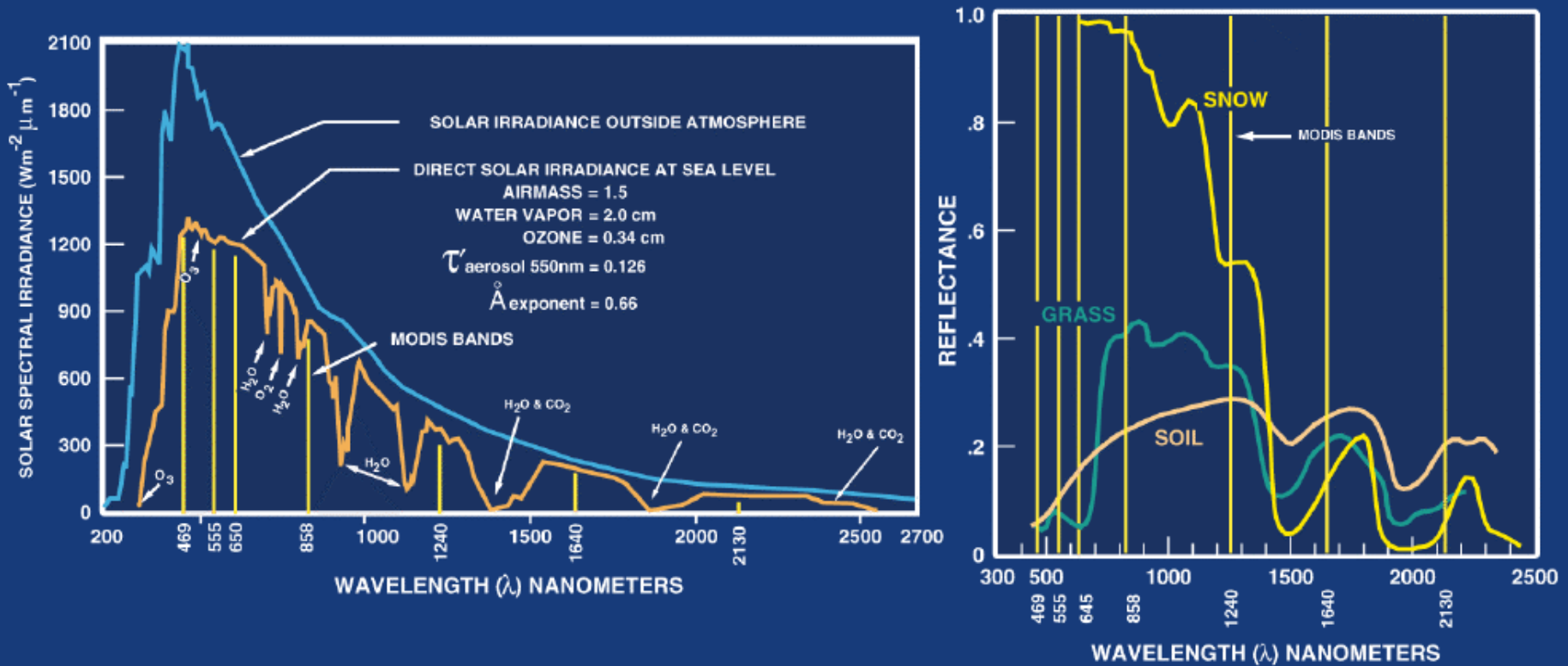


direct + diffuse

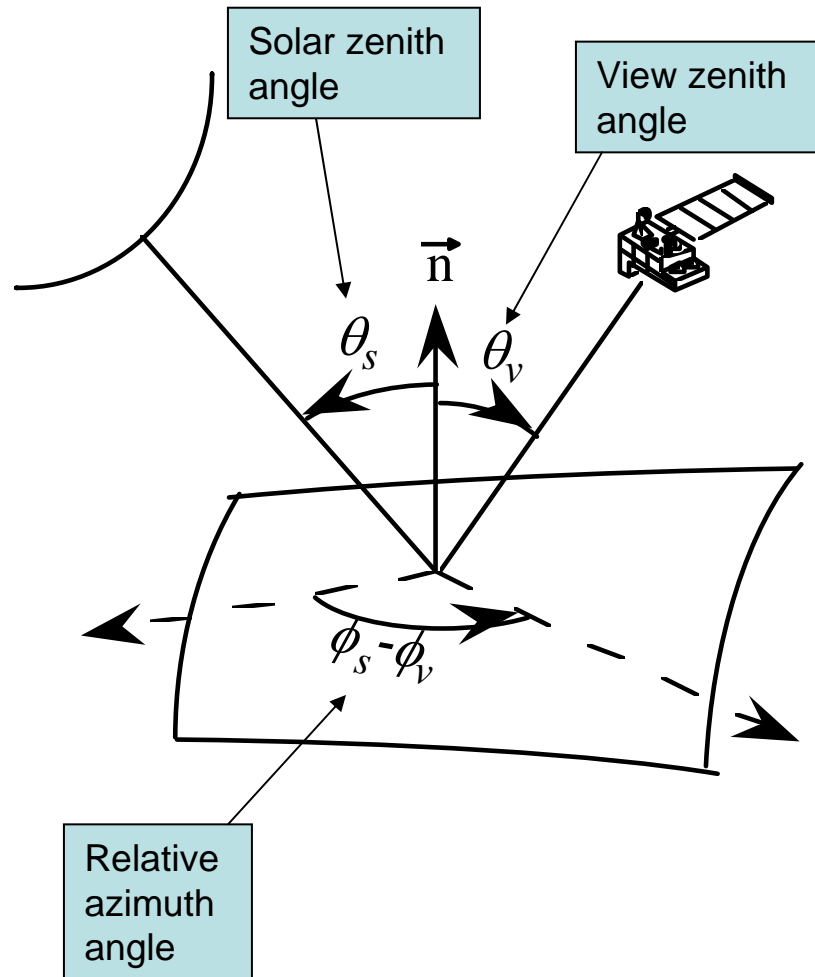


multiple scattering

Solar (reflective) spectral domain



Observation Geometry



Solution of the Radiative Transfer in the reflective domain for non absorbing atmosphere and lambertian ground

Atmospheric reflectance **Ground reflectance**
(= albedo for lambertian)

$$\rho_{app}(\theta_s, \theta_v, \phi) = \rho_{atm}(\theta_s, \theta_v, \phi) + T_{atm}(\theta_s)T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - \rho_{ground}S_{atm}}$$

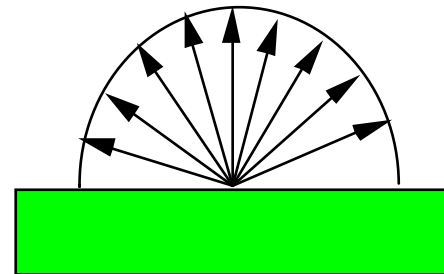
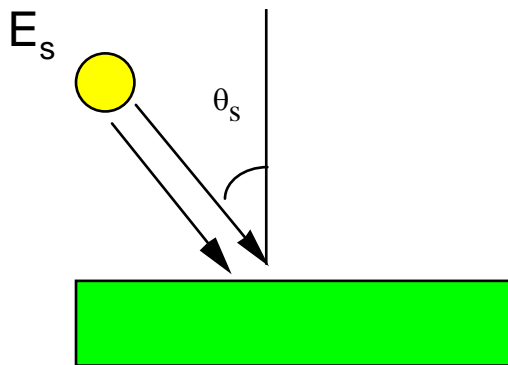
Apparent reflectance at satellite level **Atmospheric Transmissions**

Atmosphere spherical albedo

Perfect Lambertian Reflector

Radiance of the Perfect Lambertian Reflector

$$\int_0^{\pi} \int_0^{2\pi} RPLF(\theta_s, \theta, \phi) \cos(\theta) \sin(\theta) d\theta d\phi = E_s \cos(\theta_s)$$

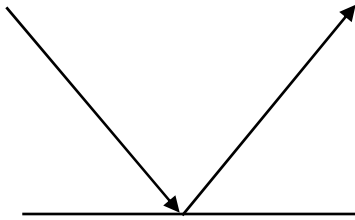


Isotropic radiation

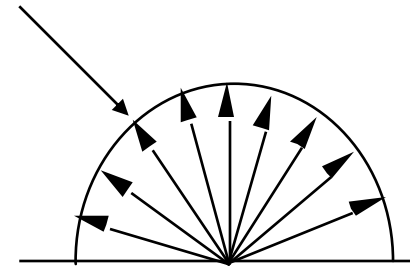
$$\rho_{\text{Perfect Lambertian reflector}}(\theta_s, \theta_v, \phi) = 1$$

$$\rho_{\text{Lambertian reflector}}(\theta_s, \theta_v, \phi) = \rho$$

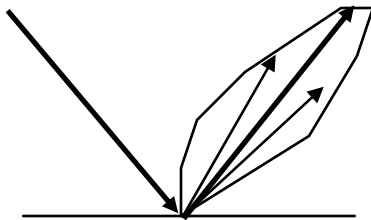
Different Types of Reflectors



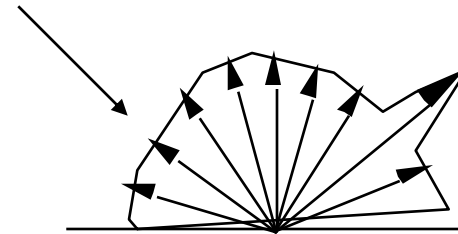
Specular reflector (mirror)



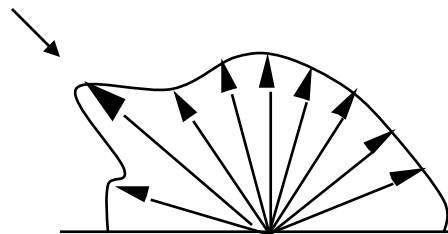
diffuse reflector (Lambertian)



Nearly Specular reflector (water)

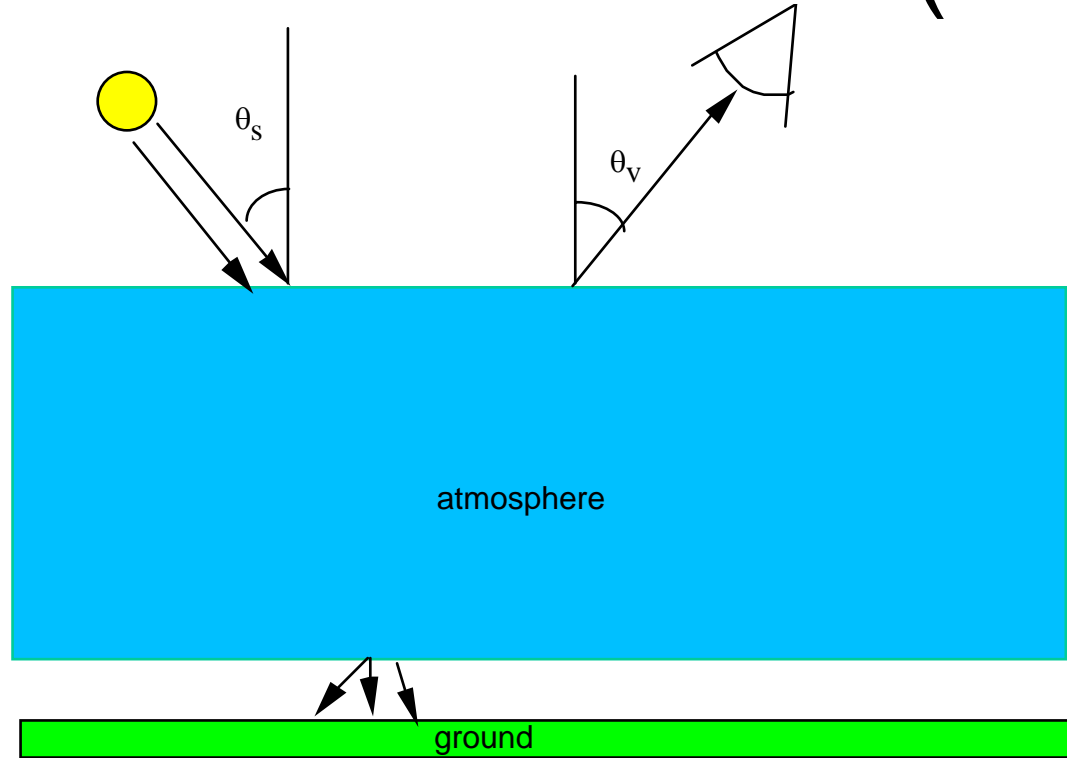


nearly diffuse reflector



Hot spot reflection

S RTE 1 interaction (cont.)

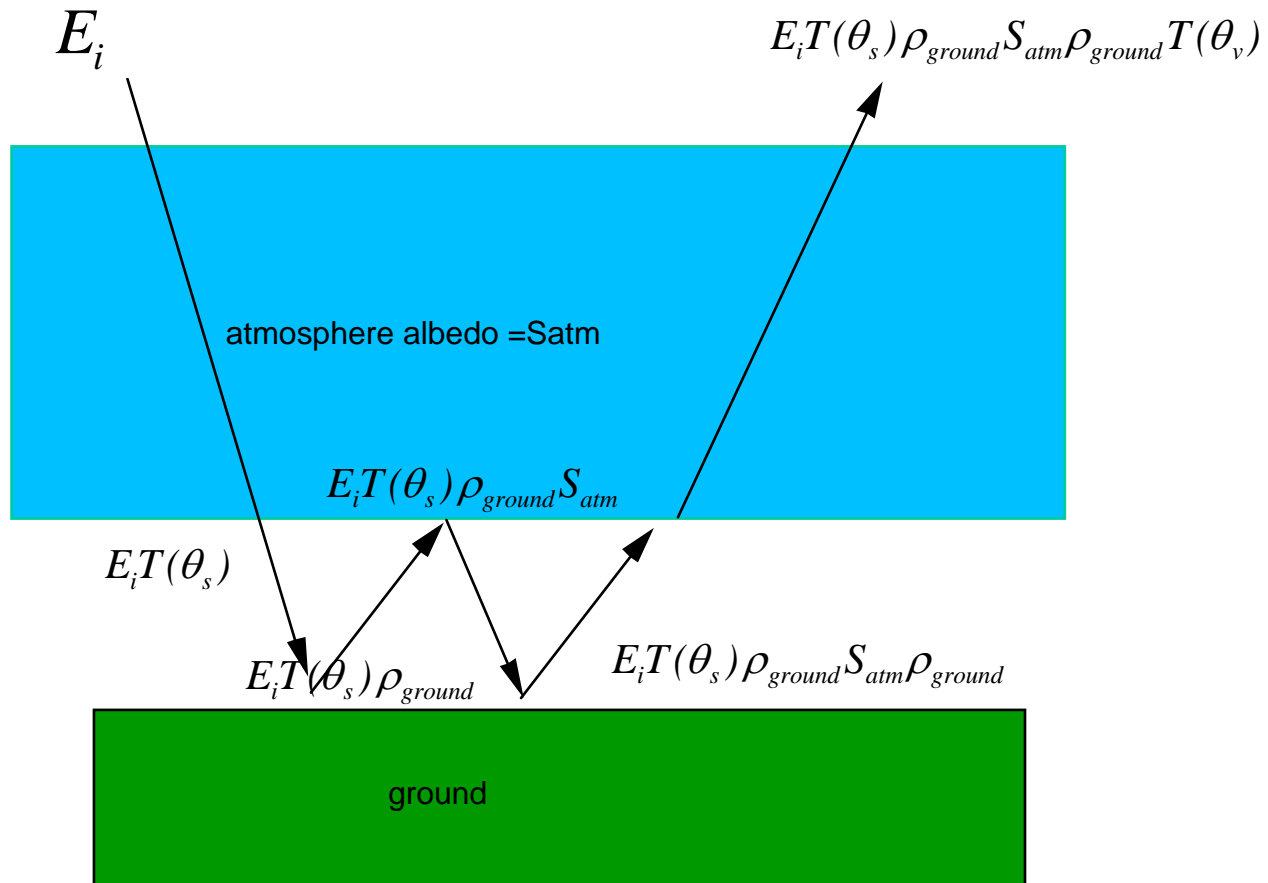


$$\rho_{app} = \rho_{atm} + \frac{E_o}{E_i}$$

$$\frac{E_o}{E_i} = \frac{T(\theta_v)E_r}{E_i} = \frac{T(\theta_v)\rho_{ground}E_t}{E_i} = T(\theta_v)\rho_{ground}T(\theta_s)$$

$$\rho_{app} = \rho_{atm} + T(\theta_v)\rho_{ground}T(\theta_s)$$

S RTE 2 interactions



S RTE Multiple Interactions

$$\rho_{app} = \rho_{atm} + T(\theta_s)T(\theta_v)\rho_{ground} \left[1 + \rho_{ground}S_{atm} + (\rho_{ground}S_{atm})^2 + (\rho_{ground}S_{atm})^3 \dots \right]$$

$$1 + r + r^2 + r^3 + \dots r^{n-1} = \frac{1 - r^n}{1 - r}$$

$\rho_{ground}S_{atm} < 1$ so when $n \rightarrow \infty$ then $(\rho_{ground}S_{atm})^n \rightarrow 0$

Therefore $\left[1 + \rho_{ground}S + (\rho_{ground}S)^2 + (\rho_{ground}S)^3 \dots \right] = \frac{1}{1 - \rho_{ground}S}$

$$\rho_{app} = \rho_{atm} + T_{atm}(\theta_s)T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - \rho_{ground}S_{atm}}$$

BRDF atmosphere coupling correction

Lambertian infinite target approximation

$$\rho_{app}(\theta_s, \theta_v, \phi) = \rho_{atm}(\theta_s, \theta_v, \phi) + T_{atm}(\theta_s)T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - \rho_{ground}S_{atm}}$$

BDRF atmosphere coupling approximation

$$\begin{aligned} \rho_{app}(\theta_s, \theta_v, \phi) &= \rho_{atm}(\theta_s, \theta_v, \phi) + e^{-\tau/\mu_s} e^{-\tau/\mu_v} \rho_s(\theta_s, \theta_v, \phi) \\ &+ e^{-\tau/\mu_v} t_d(\theta_s) \bar{\rho}_s + e^{-\tau/\mu_s} t_d(\theta_v) \bar{\rho}'_s + t_d(\theta_v) t_d(\theta_s) \bar{\bar{\rho}}_s \\ &+ \frac{T_{atm}(\theta_s) T_{R+A}(\theta_v) S_{atm}(\bar{\bar{\rho}}_s)^2}{1 - S_{atm} \bar{\bar{\rho}}_s} \end{aligned}$$

$$\bar{\rho}_s(\mu_s, \mu_v, \phi) = \frac{\int_0^{2\pi} \int_0^1 \mu L_{atm}^\downarrow(\mu_s, \mu, \phi') \rho_s(\mu, \mu_v, \phi' - \phi) d\mu d\phi'}{\int_0^{2\pi} \int_0^1 \mu L_{atm}^\downarrow(\mu_s, \mu, \phi') d\mu d\phi'}$$

$$\bar{\bar{\rho}}_s(\mu_s, \mu_v, \phi) = \overline{\bar{\rho}'_s(\mu_s, \mu_v, \phi)}$$

$$\bar{\rho}'_s(\mu_s, \mu_v, \phi) = \bar{\rho}_s(\mu_s, \mu_v, \phi)$$

STRE for non absorbing atmosphere and lambertian ground

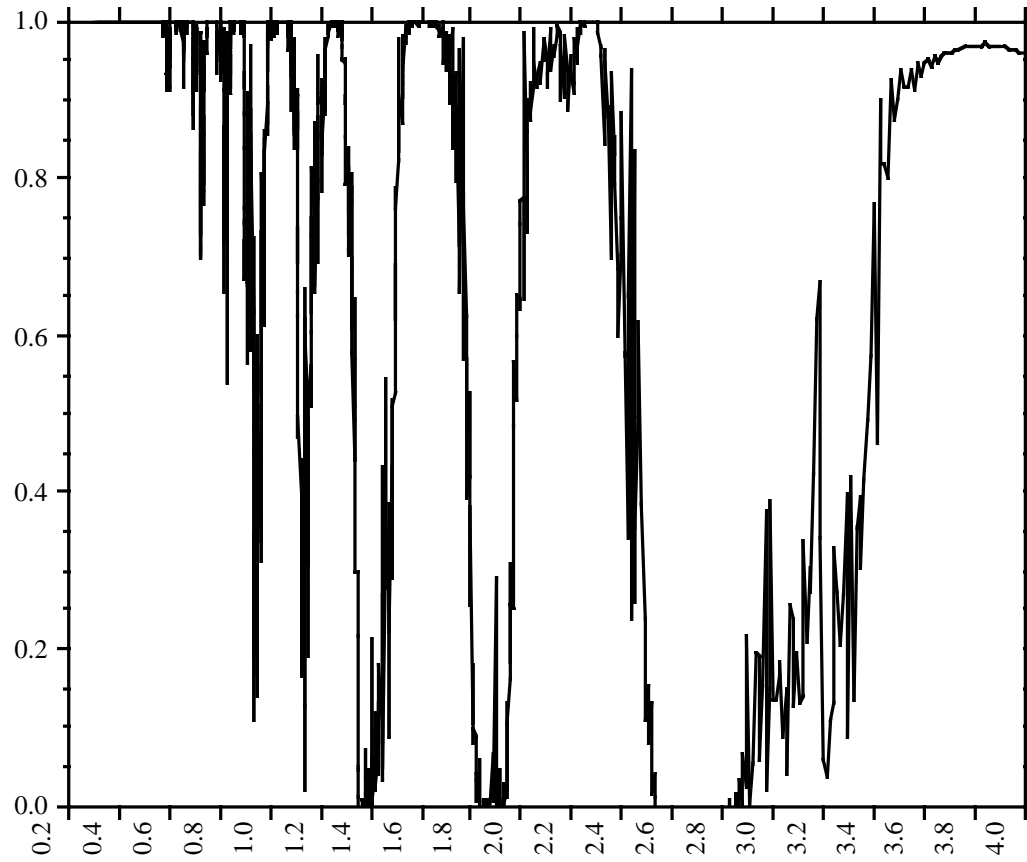
$$\rho_{app}(\theta_s, \theta_v, \phi) = \rho_{atm}(\theta_s, \theta_v, \phi) + T_{atm}(\theta_s)T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - \rho_{ground}S_{atm}}$$

Atmospheric reflectance (points to $\rho_{atm}(\theta_s, \theta_v, \phi)$)
Ground reflectance (= albedo for lambertian) (points to ρ_{ground})
Atmospheric Transmissions (points to $T_{atm}(\theta_s)T_{atm}(\theta_v)$)
Apparent reflectance at satellite level (points to $\rho_{app}(\theta_s, \theta_v, \phi)$)
Atmosphere spherical albedo (points to S_{atm})

The composition of the atmosphere

| Permanent Constituents | | Variable constituents | |
|-----------------------------------|--------------------------|--|----------------------------|
| Constituent | % by volume | Constituent | % by volume |
| Nitrogen (N ₂) | 78.084 | Water Vapor (H ₂ O) | 0.04 |
| Oxygen (O ₂) | 20.948 | Ozone (O ₃) | 12 x 10 ⁻⁴ |
| Argon (Ar) | 0.934 | Sulfur dioxide (SO ₂) ^b | 0.001 x 10 ⁻⁴ |
| Carbon dioxide (CO ₂) | 0.033 | Nitrogen dioxide (NO ₂) | 0.001 x 10 ⁻⁴ |
| Neon (Ne) | 18.18 x 10 ⁻⁴ | Ammonia (NH ₃) | 0.001 x 10 ⁻⁴ |
| Helium (He) | 5.24 x 10 ⁻⁴ | Nitric oxide (NO) | 0.0005 x 10 ⁻⁴ |
| Krypton (Kr) | 1.14 x 10 ⁻⁴ | Hydrogen sulfide (H ₂ S) | 0.00005 x 10 ⁻⁴ |
| Xenon (Xe) | 0.089 x 10 ⁻⁴ | Nitric acid vapor | trace |
| Hydrogen (H ₂) | 0.5 x 10 ⁻⁴ | | |
| Methane (CH ₄) | 1.5 x 10 ⁻⁴ | | |
| Nitrous Oxide (N ₂ O) | 0.27 x 10 ⁻⁴ | | |
| Carbon Monoxide (CO) | 0.19 x 10 ⁻⁴ | | |

Gaseous Absorption (H₂O)



Modified SRTM to account for absorption

In case of a pure molecular atmosphere (no aerosol) we can write:

$$\rho_{app}(\theta_s, \theta_v, \phi) = Tg^{other\ gases}(m, U_{gaz}) \left[\rho_{atm}(\theta_s, \theta_v, \phi) + Tg^{H_2O}(m, U_{H_2O}) T_{atm}(\theta_s) T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - S_{atm} \rho_{ground}} \right]$$

m is the air mass = $1/\cos(\theta_s) + 1/\cos(\theta_v)$

U_{gaz} is the gaz concentration

Phase function

- The phase function, $P(\Theta)$, describe the distribution of scattered radiation for one or an set of particles. It is normalized

such as:

$$\int_0^{2\pi} \int_0^{\pi} P(\Theta) d\omega = 4\pi$$

since

$$\int_0^{2\pi} \int_0^{\pi} P(\Theta) \sin(\theta) d\theta d\phi = 2\pi \int_0^{\pi} P(\theta) \sin(\theta) d\theta$$

we have

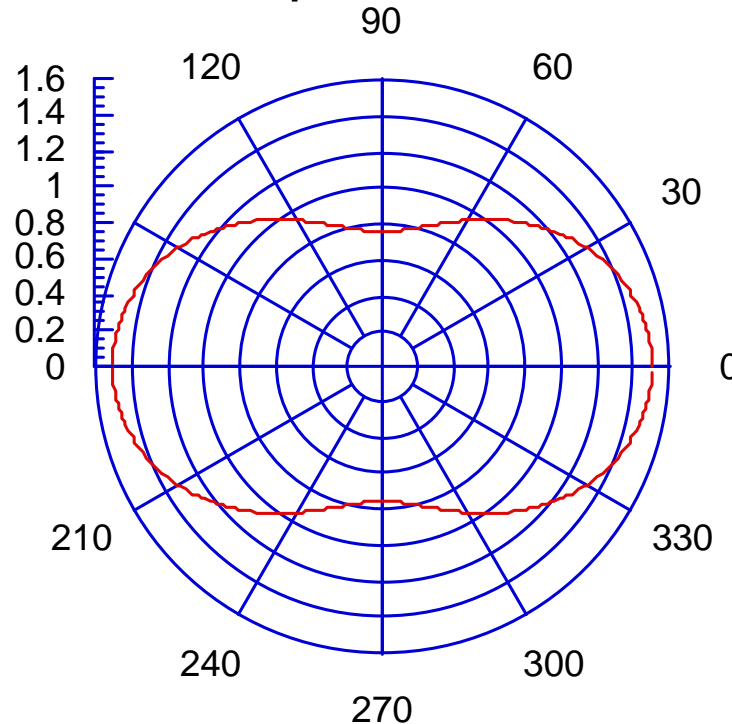
$$\int_0^{\pi} P(\theta) \sin(\theta) d\theta = 2$$

Rayleigh/molecular scattering

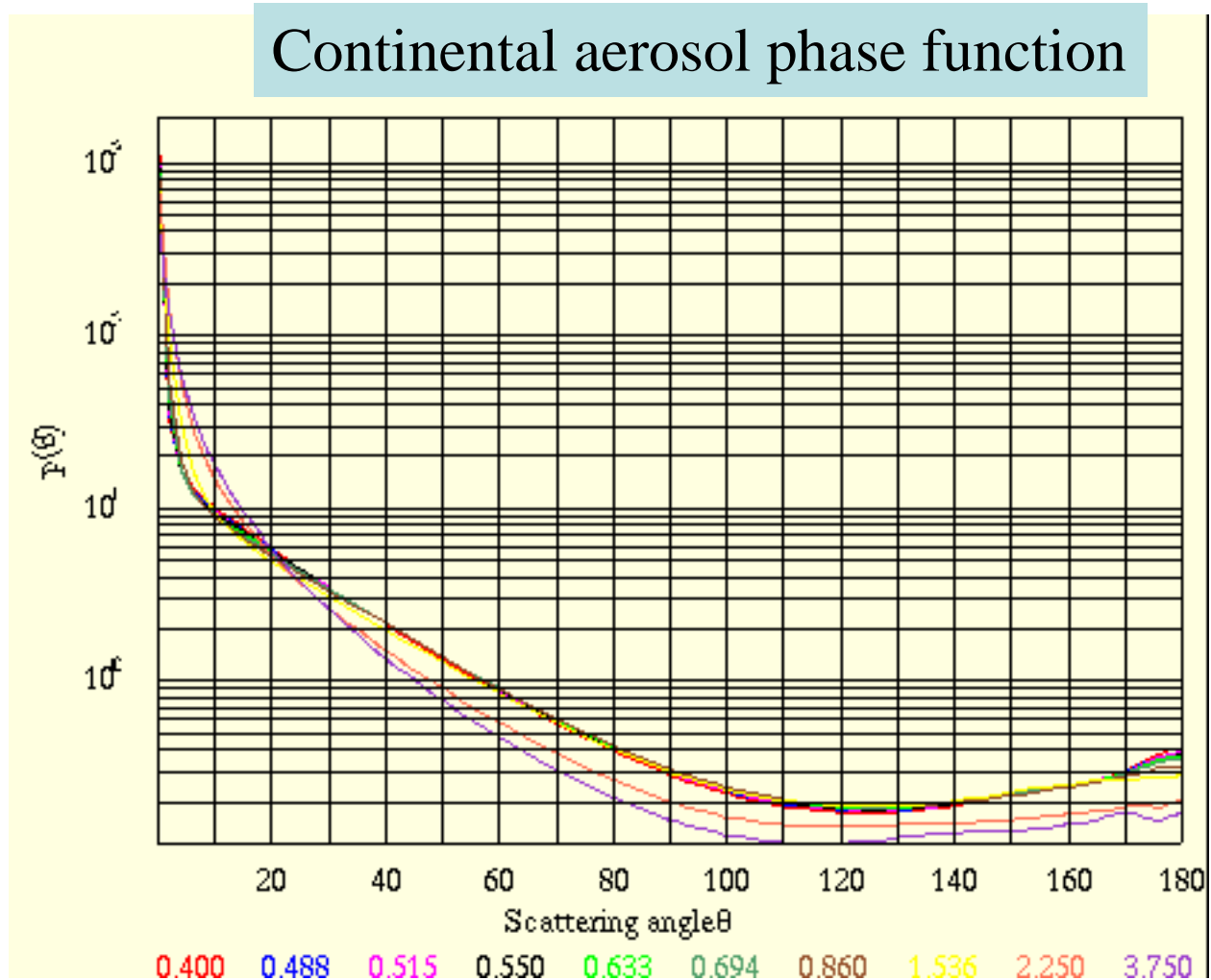
1/4

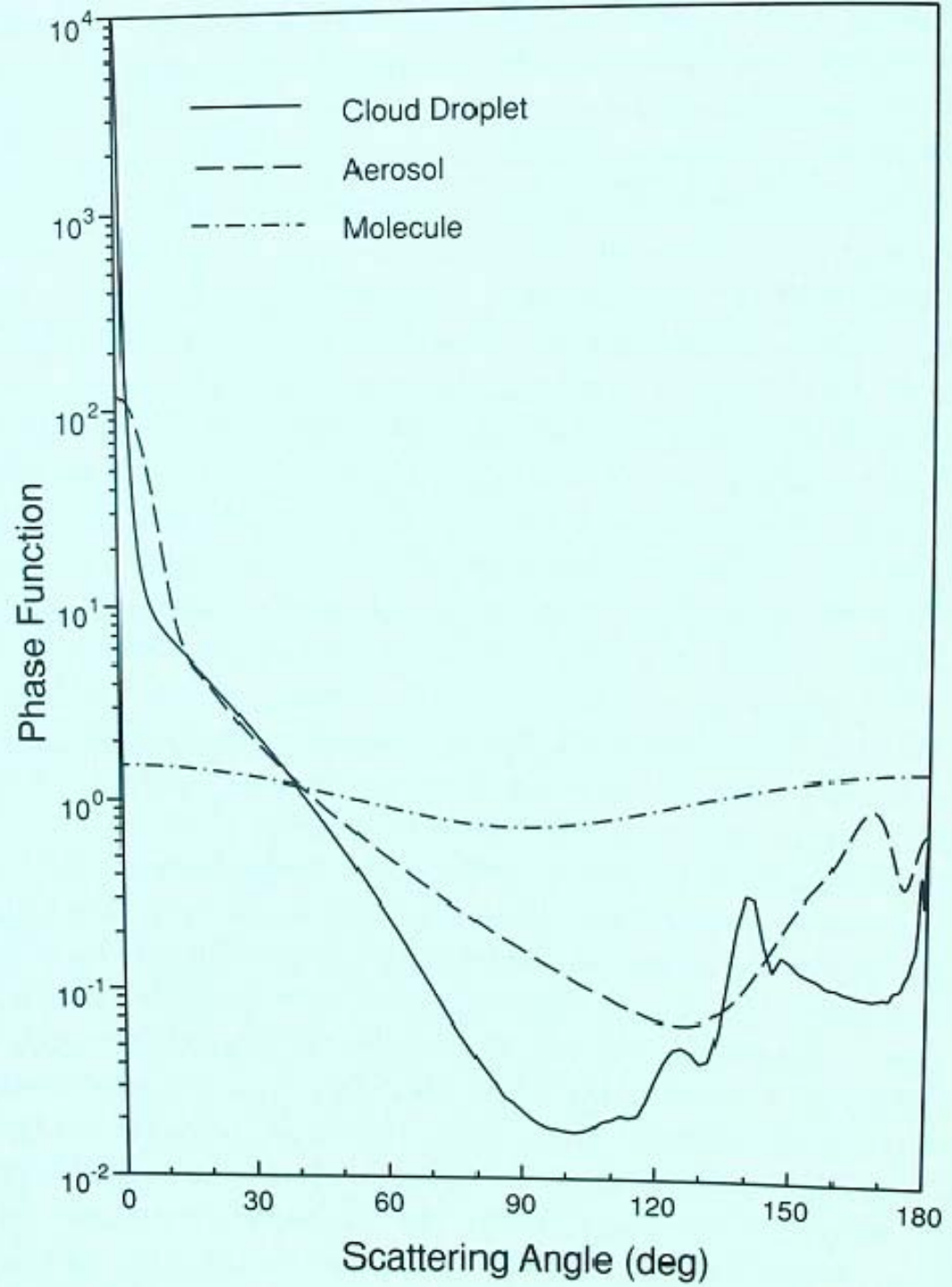
- Rayleigh or molecular scattering refers to scattering by atmospheric gases, in

that case:

$$P(\Theta) = \frac{3}{4} (1 + \cos^2(\Theta))$$


Aerosol scattering 2/5





Some Useful Radiative Transfer codes:

Mie codes:

Homogeneous Sphere:

ftp://climate1.gsfc.nasa.gov/wiscombe/Single_Scatt/Homogen_Sphere/Exact_Mie

Layered Sphere:

ftp://climate1.gsfc.nasa.gov/wiscombe/Single_Scatt/Coated_Sphere

Disort:

For Fortran 77:

ftp://climate1.gsfc.nasa.gov/wiscombe/Multiple_Scatt/DISORT1.2

For Fortran 90:

ftp://climate1.gsfc.nasa.gov/wiscombe/Multiple_Scatt/DISORT1.3

Other Multiple Scattering Codes:

ftp://climate1.gsfc.nasa.gov/wiscombe/Multiple_Scatt

SBDART:

<http://www.paulschou.com/tools/sbdart/>