

1. To determine the volumetric flow rate, one needs to use the following equation:

$$Q = \frac{K \Delta H m}{n}$$

where K is hydraulic conductivity, ΔH is the total hydraulic head difference, m is the number of flow tubes and n the number of equipotential intervals. At the top of the soil, the pressure head $h = 100$ cm = 1 m. At the drain, the pressure head is 0. The elevations at the top and the drain are 100 m and 50 m, respectively. Therefore, we have the hydraulic heads at the top and drain of 101 m and 50 m, respectively. We thus have $\Delta H = 101 - 50 = 51$ m. From the flow net, we count 5 equipotential intervals and 6 flow tubes, or $n = 5$ and $m = 6$. We thus have the volume of water collected per hour as

```
> K := 5e-6*60*60; DeltaH := 101.0 - 50.0; m := 6.0; n := 5.0; dQ
:= K*DeltaH/n; Q := K*DeltaH*m/n;
      K := 0.018000
      DeltaH := 51.0
      m := 6.0
      n := 5.0
      dQ := 0.1836000000
      Q := 1.1016000000
```

2. To determine the storage coefficient, one needs to calculate the water capacity of the soil. The water capacity at 0 cm is zero, because when $h = 0 > h_d$ the water content is always the same as θ_s . For $h = -10$ cm and -1000 cm, the corresponding water capacity values are:

```
> theta[r] := 0.03; theta[s] := 0.41; N := 0.75; m := 4.25; hd :=
-4.0; h := -10.0; WC10 := -N*(theta[s]-theta[r])/h*(hd/h)^N;
      theta_r := 0.03
      theta_s := 0.41
      N := 0.75
      m := 4.25
      hd := -4.0
      h := -10.0
      WC10 := 0.01433474110
> h := -1000; WC1000 := -N*(theta[s]-theta[r])/h*(hd/h)^N;
      h := -1000
      WC1000 := 0.4533043155 10^-5
```

With the given water and soil matrix compressibility, one has the storage coefficient at $h = 0$ cm as:

```
> rho[w] := 1000; g := 9.81; beta[w] := 4.8e-10; beta[p] :=
```

```

6.9e-8;S[s0] := rho[w]*g*(beta[p]*1.0+beta[w]*theta[s])+0.0;
          rho_w := 1000
          g := 9.81
          beta_w := 0.48 10^-9
          beta_p := 0.69 10^-7
          S_s0 := 0.0006788206080

```

At $h = -10$ cm, we have $\frac{d\theta}{dh} = WC10$ as calculated above, therefore:

```

> h := -10; theta[10] := theta[r] + (theta[s] -
theta[r])*(hd/h)^N; S[s10] :=
rho[w]*g*(beta[p]*theta[10]/theta[s]+beta[w]*theta[s])+WC10*100
;
          h := -10
          theta_10 := 0.2211298813
          S_s10 := 1.433841115

```

At $h = -1000$ cm, we have $\frac{d\theta}{dh} = WC1000$ as calculated above, therefore:

```

> h := -1000; theta[1000] := theta[r] + (theta[s] -
theta[r])*(hd/h)^N; S[s1000] :=
rho[w]*g*(beta[p]*theta[1000]/theta[s]+beta[w]*theta[s])+WC1000
*100;
          h := -1000
          theta_1000 := 0.03604405754
          S_s1000 := 0.0005147419042

```

When the soil column is unsaturated and if we use the upward z as the positive z direction, the

hydraulic gradient $\frac{dH}{dz} = 1$, given that the pressure head is uniformly at -10 cm over the length of the soil column. We may thus calculate the hydraulic conductivity, specific discharge and volumetric flow rate as follows:

```

> h := -10; K[s] := 0.24; K[10] := K[s]*(hd/h)^m; q :=
-K[10]*(1.0); r := 5.0; A := 3.14159*r^2; Q := q*A;
          h := -10
          K_s := 0.24
          K_10 := 0.004886143358
          q := -0.004886143358
          r := 5.0
          A := 78.5397500

```

$$Q := -0.3837564778$$

with the negative signs of q and Q indicating the direction of the flow or downwards.

- Using the tangent defraction law, the relationship between hydraulic conductivity and defraction angle is:

$$\frac{K_1}{K_2} = \frac{\tan(\theta_1)}{\tan(\theta_2)}$$

Given the following values for the top clay layer:

```
> K[1] := 5e-6; theta[1] := 5.0;
      K1 := 0.5 10-5
      θ1 := 5.0
```

one may calculate the angle between the interface of the clay and sand layers, or θ_{clay} as:

```
> K[2] := 4e-9; theta[2] :=
  arctan(K[2]/K[1]*tan(theta[1]*3.14159/180.0))/3.14159*180;
  theta[clay] := 90.0 - theta[2];
      K2 := 0.4 10-8
      θ2 := 0.004010184917
      θclay := 89.99598982
```

The water enter the bottom sand layer with angle as calculated in the following:

```
> theta[3] :=
  arctan(K[1]/K[2]*tan(theta[2]*3.14159/180.0))/3.14159*180;
  theta[sand] := 90.0 - theta[3];
      θ3 := 5.000000001
      θsand := 85.00000000
```

which suggests that the angles of entering and existing the clay layer are the same. The flow lines and equipotential lines may be plotted as follows:

